2.	From	Error to S	Signal
To build a model that is complecation calibrated measurements, and seems reasonable to characted developing their numerical model model <i>t</i> . Thus, when we representation error from <i>C</i> and	ete and nonli <i>t</i> , the simples rize geophysi lels. Similarly, equate meas a statistician	inear from the outs st signal model (ign icists as attentive to statisticians are at sures and model, might add equation	et, we begin b noring noise for to unresolved s tentive to limita a geophysicis error to <i>t</i> :
Calibrated • $C - \epsilon_C^{REP}$	=	$t + \epsilon_C^{EQ}$	<b>Equ</b> accomm signal ir
Uncalibrated • $U - \epsilon_U^{REP}$	$= \alpha_U +$	$\beta_U t + \epsilon_U^{EQ}$	betweer
Measure - representation error		Model + model erro	r Repres accom in C or
Ce t cen.wiki	eci n'est pas un éléphant ipedia.org/wiki/Asian_elephant) All models are wrong	Measurement Error Models WAYNE A. FULLER	Althoug errors se need to a geop statisticia with t
In the literature, it is rare to find the uncalibrated measurement $e$ U is sometimes seen as a debat to accommodate representation error in $C$ and $U$ as well.	a separate e equation. An a te. Perhaps v error in C ar	quation error or rep accommodation of r ve are just "ripping on nd <i>U</i> , then we migh	resentation error representation error off the bandaid" t need to accor
Including a separate equation e and $U$ are equally justified, even be no other way for this model only measurements are $C$ and justification that is more practical	error in <i>U</i> can n though their to incorporate d <i>U</i> , and the al, as it relates	be justified by its line r numerical values of the impact of hidde se may be impacted to the model solution	near relationshi differ). Moreove en or confoundi ed differently. on.)
)n 17. aug. 2017 10:49, GRAHA Hi Rick	M DUNN wrote:	Equation e	rror can be <b>co</b> and <b>co</b>
I had wondered whether your rather than measurement erro don't think it has any pract interpretation). I can't see not be correlated.	correlated error ors (or a combina tical implication e any reason why	s were equation errors ation of both) but I is (just a different equation errors should	with oth
With best wishes Graham			Nonl in all n

Graham Dunn Professor of Biomedical Statistics Centre for Biostatistics University of Manchester

In any case, the interpretation of error in terms of signal is challenging, but in 2017, Graham Dunn provided a five-line comment. This took some time to absorb, but basically, he said that equation error could be mixed in with other errors and equation error could be correlated.

If we carry this advice forward, and place measurements on the LHS and model terms on the RHS, then we have linear association, correlated and uncorrelated equation error, and representation error, which is uncorrelated by definition. All are signal components insofar as they are needed to describe C and U even without measurement error. Without loss of generality, measurement error in C and U can be divided among the signal terms on the RHS, or written as its own term (as we wish):

$\sim$		EO DED	Measures = linear association
C	=	$t + \epsilon_C^{L_{\mathcal{Q}}} + \epsilon_C^{L_{L_{\mathcal{P}}}}$	+ correlated equation
U	=	$\alpha_U + \beta_U t + \epsilon_U^{EQ} + \epsilon_U^{REP}$	+ uncorrelated equation and representation + measurement error (in

Since we are interested in terms that can be evaluated numerically, we follow Prof. Dunn's advice and let correlated equation error stand alone and combine uncorrelated equation error and representation error into a separate term:

C = $t + \epsilon + \epsilon_C$ Measures = linear association + nonlinear association + lack of association  $U = \alpha_U + \beta_U t + \epsilon + \epsilon_U$ 

The core of our wavelike measurement model is here. This form makes explicit a "systematic error" term  $\varepsilon$  that we would call error cross-correlation, but whose genuine interpretation is nonlinear association. Similarly, our "random error" terms have the *genuine* interpretation of a lack of association. Although signal defines our interpretation of terms, again, we can conceive of spurious components in each term (https://arxiv.org/abs/2110.08969 is a more formal derivation).

by equating C, the r the moment). It scales in C when ations of the linear ist might subtract

#### uation error

nodates nonlinear n the relationship en C and U (via t).

esentation error nmodates scales U, but not *both*.

ah we write these eparately, we also ask to what extent physicists and ans are concerned the same error.

or term included in error in both C and ', in that if we want mmodate equation

ip to t (note that C ver, there seems to ling variables. The (There is a third

orrelated ombined ner errors

linearity measures

Is that still called linear regression?

> Signal error error all terms) Noise

### A Framework for Exploring Systematic Error in a Quantitatively Complete Measurement Model

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Can systematic and random error, as defined in VIM (2008), also be interpreted in a way that follows Pearson's (1902) emphasis of "genuine error"? Respectively, we submit that "nonlinear association" and "lack of association" are their genuine interpretations. First, we develop these ideas out of a purely statistical notion of a "measurement model", or regression model. Then, we examine a playful controlled experiment in the hydrological context.

## 1. Sampling Signal as a Wave

Our wavelike measurement model employs a collocated sampling strategy that is denoted by the coloured dots (moving left to right). Of interest is the linear, nonlinear, and lack of association between a less familiar "uncalibrated" platform or instrument, and one that is more familiar, like an in situ platform. It is notable that wavelike sampling is possible for large datasets that are gridded in space or time; not every scientific discipline has the modelling freedom that comes with such extensive sampling.

Cal/val phase of Uncalibrated platform
STUVW
linear nonlinear lack of
ABCDE
Calibrated platforms aliders / moorings / satellit

We employ a trichotomy (e.g., truth + systematic + random error) to describe individual model terms. Model heritage can be traced to an introduction by Pearson (1902) of three observers as a means to identify the errors of each. The geophysical focus on Pearson's approach re-emerged as triple collocation in 1988, and helped motivate our wavelike model. However, instead of three independent observers (or three sets of red dots), method-of-moments solutions are identifiable using "predictive samples" for just two datasets. Whereas a collocation of three or more datasets is rare, our wavelike model still requires successive sampling of two datasets:

A B C D E

	• S	=	$\alpha_S + \beta_S t + \lambda_S$	$(\lambda_T (\epsilon + \epsilon_U))$
	• T	=	$\alpha_T + \beta_T t +$	$\lambda_T \left( \epsilon + \epsilon_U \right)$
	• U	=	$\alpha_U + \beta_U t +$	$\epsilon + \epsilon_U$
	• V	=	$\alpha_V + \beta_V t +$	$\lambda_V \left( \epsilon + \epsilon_U \right)$
	• W	=	$\alpha_W + \beta_W t + \lambda_W$	$\epsilon \left( \lambda_V \left( \epsilon + \epsilon_U \right) \right)$
1 km				
1	• A	=	$\alpha_A + \beta_A t + \lambda_A$	$(\lambda_B (\epsilon + \epsilon_C))$
• or 1 day	• B	=	$\alpha_B + \beta_B t +$	$\lambda_B \left( \epsilon + \epsilon_C \right)$
UTUdy	• C	=	t +	$\epsilon + \epsilon_C$
	• D	=	$\alpha_D + \beta_D t +$	$\lambda_D \left( \epsilon + \epsilon_C \right)$
	• E	=	$\alpha_E + \beta_E t + \lambda_E$	$(\lambda_D (\epsilon + \epsilon_C))$

Here,  $\alpha$ ,  $\beta$ , and t capture linear association between samples, with symmetric first-order autoregressive errors for each dataset separately. Errors are also correlated in  $\varepsilon$ , which is a term not always included in canonical models (Fuller 2006). Geophysical theory (equations of motion) allows for a nonlinear relationship between any two samples (like *B-C* or *C-U*). The key question is how a wavelike model accommodates nonlinearity in what looks to be a set of linear relationships?

#### 4. Recommendations and References

Progress or lack of progress (challenges) so far Mahalanobis (1947) seems to address the metrological practice of expanded uncertainty, but is unclear about its motivation. Meteorologists might pick up where Mahalanobis (1947) left off [less with reference to Type-A/B cyclogenesis (Bosart 1994) and more with reference to what Stigler (2018) refers to "high serial correlations" in atmospheric soundings]. Pearson (1902) and Mahalanobis (1947) seem to be advocating for genuine interpretations, but equation error (Fuller 2006) only develops later.

(2) Indication of what is needed to improve the information content in environmental measurements or models Metrologists might distinguish quantitative models to address a) "what is the process?" and b) "how do we measure the process?" Even for the same process, these are not necessarily the same model. We are familiar with a hierarchy of process models (Held 2015), but perhaps not with a hierarchy of measurement (or regression) models. One tenet of statistical inquiry is a model with two terms: observation = truth + error (Salsburg 2017), but metrological convention admits a further division into systematic and random error. We submit that the corresponding nonlinear model with three terms (observation = linear + nonlinear + unassociated) also has quantitative solutions.

(3) Proposals of new or existing tools not currently in use, for improving information content Shameless self-promotion: "numerical consensus" methods to solve this nonlinear model are at https://github.com/JuliaAtmosOceanHydro/MeasurementModelDemos

(4) Description of potential means to achieve that improvement One may appeal to colleagues who wish to learn about nonlinear processes that, at the moment, they might be training a neural network to emulate. (Outside our fields,

some colleagues also deal with nonlinear processes that are downright spooky.) (5) Opportunities for collaboration between communities to achieve these improvements

There seems to be an opportunity to bring statisticians and data scientists closer together (Donoho 2017)



gliders / moorings / satellites / numerical models

```
(\epsilon) + \epsilon_T ) + \epsilon_S
() + \epsilon_T
(+) + \epsilon_V
(\cdot) + \epsilon_V ) + \epsilon_W 
(\epsilon) + \epsilon_B ) + \epsilon_A
(+) + \epsilon_B
```

S, T, V, W are autocorrelated *instruments* (proxies) of U

A, B, D, E are autocorrelated *instruments* (proxies) of *C* 

 $+\epsilon_D$  $(\epsilon) + \epsilon_D ) + \epsilon_E$ 

or "predictive samples" ("forecasts" or "revcasts")

References are provided a https://docs.google.com/document/d/1ukxYVb0Ec8DERQIzqh99DiByYRM97QEPES5sniROTiM

# 3. A Control Experiment

Returning to Prof. Dunn's question of practical implications, there is no doubt about a different interpretation: measurements, variance, and covariance all have a nonlinear component, and thus, Pearson correlation also has a nonlinear component. Such interpretations can be quantified. By random samples of uniform and Gaussian distributions, we can simulate anomalous river height or river slope at 10000 well separated reaches that have been observed hourly since the turn of the year 1900. The anomaly is such that there is only variability between a day and a year:

- For each river, signal is slow and
- wavelike (daily-to-yearly)10000 well-separated rivers sampled
- (STUVW and ABCDE) at 1-h intervals • Perfect calibration :  $\beta_{T} = 1$ (STUVW = ABCDE)

For each river reach, we generate about a million samples of a uniform random distribution, bandpass filter to get a timeseries in the range -0.2 to 0.2, and from the middle of each timeseries (like the one above), we take 5 consecutive samples to define the slowly varying signal in both C and U, which are perfectly calibrated. This experiment takes up the challenge of trying to simulate a confounding impact in the presence of measurement error, so two types of perturbations are employed, and both are designed to be resolved by a lack of association between C and U:





Gaussian perturbations are added to U in varying degrees. Spearman perturbations are a random reordering of signal in U that might appear to treat two river reaches as exchangeable, but are meant to simulate a confounding impact without imposing a systematic relationship between C and U. Although Gaussian and Spearman perturbations might appear to be similar in the upper and lower panels, by design, Spearman perturbations impose no turbulent variations across STUVW. Spearman perturbations can be said to perturb U alone, but because they are a perturbation of relative order, it follows that if we consider C to be unperturbed and equal to the baseline signal, and U alone to be perturbed, then it is only predictive samples involving STUVW, but not ABCDE, that access both Gaussian and Spearman perturbations, and hence, might distinguish them on this basis.



Model solutions by the method of moments require only six sample sets, ABCDEU or CSTUVW. Each set yields 6 variance equations, 15 covariance equations, and 17 unknowns. Analytic solutions exist for 15 of these unknowns, but the remaining two unknowns are the variance of t and  $\beta_{ij}$ , the two key parameters. We minimize the distance of a solution to these two parameters using minima or pathways in each of six covariance equations that involve the variables other than U and C. Together with the constraint that variance is positive, as expected, we refer to this equitable accommodation of weak or imperfect constraints as an exercise in consensus building.

Wavelike Solutions	S •	Т •	U •	V	И •
			_		

- Noisy subsamples
- Consensus solutions
- Better bounds on  $\beta_{II}$

Our two wavelike solutions happen to be directly comparable to ordinary and reverse regression solutions that are analytic. So, in the case of our wavelike solutions, STVW and ABDE, they provide sub at allow us to estimate Gaussian error, and in the case of analytic l ordinary regression, we would assume that those samples s T U V W • • • • • e is no error in the U or C variable bracketed by them.

subsamp	les	arou	nd	U	and	С	tha
U	SO	lution	s b	y r	ever	se	and
$\begin{array}{c}\bullet\\\bullet\\\bullet\\\bullet\\\bullet\\\bullet\\\bullet\\\bullet\\\bullet\\\bullet\\\bullet\\\bullet\\\bullet\\\bullet\\\bullet\\\bullet\\\bullet\\\bullet\\\bullet\\$	dc	n't ex	kist	an	d tha	at th	nere

																				U	
$\sigma_t^2$	$\beta_U~({ m O}/{ m WAV}/{ m R})$	$\operatorname{Var}(C)$	LA	NA	UR	$\operatorname{Var}(U)$	LA	NA	UR	Gauss	Spearman	$\sigma_t^2$	$\beta_U~({ m O}/{ m WAV}/{ m R})$	$\operatorname{Var}(C)$	LA	NA	UR	$\operatorname{Var}(U)$	LA	NA	UR
6.46	0.95/0.96/1.08	6.57	98.25	0.60	1.15	6.75	87.88	0.58	11.54	/	~	5.79	0.95/1.04/1.08	6.57	88.13	3.28	8.59	6.75	92.47	3.20	4.33
6.47	0.95/0.96/1.21	6.57	98.50	0.31	1.20	7.52	79.25	0.27	20.48			5.30	0.95/1.04/1.21	6.57	80.71	11.22	8.07	7.52	75.68	9.81	14.51
6.44	0.95/0.96/1.41	6.57	97.95	0.68	1.37	8.79	67.63	0.51	31.86	1		4.88	0.95/1.03/1.41	6.57	74.26	18.41	7.32	8.79	58.78	13.76	27.46
6.45	0.95/0.96/1.70	6.57	98.14	0.41	1.45	10.57	56.46	0.25	43.29	1		4.42	0.95/1.02/1.70	6.57	67.23	26.30	6.47	10.57	43.44	16.34	40.22
6.43	0.95/0.96/2.06	6.57	97.80	0.60	1.60	12.85	46.39	0.31	53.31	1		3.94	0.95/1.01/2.06	6.57	60.00	34.18	5.82	12.85	31.36	17.44	51.20
6.46	0.85/0.86/1.21	6.57	98.36	0.35	1.28	6.75	70.75	0.34	28.90	1	1	4.68	0.85/1.16/1.21	6.57	71.24	2.57	26.20	6.75	92.67	2.50	4.84
6.47	0.85/0.86/1.35	6.57	98.50	0.20	1.30	7.52	63.75	0.18	36.07		1	4.30	0.85/1.13/1.35	6.57	65.45	10.69	23.86	7.52	73.75	9.34	16.91
6.46	0.85/0.86/1.57	6.57	98.35	0.27	1.38	8.79	54.56	0.20	45.24	1	1	3.88	0.85/1.11/1.57	6.57	59.09	19.23	21.68	8.79	55.00	14.38	30.62
6.43	0.85/0.86/1.89	6.57	97.87	0.57	1.56	10.57	45.30	0.36	54.34	1	1	3.55	0.85/1.13/1.89	6.57	54.06	23.80	22.14	10.57	43.15	14.79	42.06
6.42	0.85/0.86/2.30	6.57	97.62	0.65	1.73	12.85	37.29	0.33	62.38	1	1	3.05	0.85/1.12/2.30	6.57	46.39	33.32	20.29	12.85	29.56	17.02	53.42
6.35	0.63/0.64/1.62	6.57	96.68	1.15	2.17	6.75	39.06	1.12	59.83	1	0	2.65	0.63/1.54/1.62	6.57	40.28	1.63	58.08	6.75	92.34	1.59	6.07
6.37	0.63/0.64/1.80	6.57	96.89	0.95	2.16	7.52	35.23	0.84	63.93		0	2.41	0.63/1.39/1.80	6.57	36.65	12.03	51.32	7.52	63.62	10.51	25.87
6.38	0.63/0.65/2.11	6.57	97.04	0.83	2.13	8.79	30.21	0.62	69.17	1	0	2.27	0.63/1.45/2.11	6.57	34.47	13.13	52.40	8.79	55.33	9.81	34.86
6.40	0.63/0.65/2.54	6.57	97.34	0.57	2.09	10.57	25.26	0.36	74.39	1	0	2.03	0.63/1.48/2.54	6.57	30.94	17.78	51.28	10.57	42.18	11.04	46.78
6.43	0.63/0.65/3.08	6.57	97.80	0.22	1.97	12.85	20.90	0.11	78.98	0	0	1.57	0.63/1.51/3.08	6.57	23.95	27.50	48.55	12.85	27.62	14.05	58.33
Our lower bound on $\beta_{U} = 1$ (WAV) is			No Spearm	distin	ction : mixed			Our upper	bound on $\beta_{U} = 1$	(WAV) is	S \	Sp	bearma proxy	an		Ga	ussian proxv				
											i levelse legie	331011 (1)	)		. ,			I.	,		

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ABCDE

Linear (LA), nonlinear (NA), and lack of association (UR) results are expressed as normalized U and C variance budgets, averaged over 100 simulations. By design, ABDE are poorly employed, because they don't sample any Gaussian perturbations, which makes this model solution perfectly consistent with ordinary regression, which assumes C is error-free. By contrast, STVW are well employed in sampling Gaussian perturbations. This model solution functions an upper bound on  $\beta_{\mu}$ , but provides a better bound than reverse regression. The lack of association terms in C and U allow us to distinguish Spearman and Gaussian perturbations, and we suspect that a model with all three

association categories is needed to do so.







